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# **"THE PROPER USE OF TROPICAL PEATLAND"**



Edited By: Suwardi, Ryusuke Hatano, Basuki Sumawinata, Darmawan, Suwido Limin, Yanetri Asi Nion



## INTEGRATED FIELD ENVIRONMENTAL SCIENCE -GLOBAL CENTER OF EXCELLENT (IFES-GCOE) INDONESIAN LIAISON OFFICE

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## Proceedings of Palangka Raya International Symposium and Workshop On Tropical Peatland Management " The Proper Use of Tropical Peatland"

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## SOIL GREENHOUSE GASES EMISSIONS FROM VARIOUS LAND USES IN TROPICAL PEATLANDS IN INDONESIA

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#### ABSTRACT

Accurate estimation of greenhouse gases (GHGs;  $CO_2$ ,  $N_2O$  and  $CH_4$ ) fluxes from tropical peatlands is important because of their high GHGs emissions. Soil GHGs fluxes from several land uses (natural forest, burnt forest, cropland, acacia plantation) in Indonesian tropical peatlands were measured monthly through one year, and the relationship between the fluxes and soil environmental variables was studied. Soil  $CO_2$  flux was higher in cropland and smaller in forest and burnt forest, intermediate in plantation. Several peaks of the  $CO_2$  flux were observed in rainy season. Soil  $N_2O$  flux also showed peaks in rainy season in all landuses, however, the  $N_2O$  flux was considerably high in cropland compared with other lands. Soil  $CH_4$ flux was higher in burnt forest than other land uses due to flooding in burnt forest during rainy season. Average water table depth correlated significantly to average  $CO_2$  flux positively, average  $N_2O$  flux exponentially, and average  $CH_4$  flux negatively. However, the contribution of  $CH_4$  emission to global warming potential (GWP) was negligible in water table above peat surface due toflooding. The GWP increased from 40 to 140t  $CO_2$  ha<sup>-1</sup> yr<sup>-1</sup> with drop of water table from 0 to 100 cm, in which  $CO_2$  emission increased from 40 to 100 t  $CO_2$  ha<sup>-1</sup> yr<sup>-1</sup>, and  $N_2O$ emission contributed to the GWP gradually after water table became deeper than 50 cm.

Keywords : Carbon dioxide, ground water table, methane, nitrous oxide, tropical peatland

#### INTRODUCTION

Tropical peatlands include huge amount of carbon (C) in the small land area, that is 70 GtC in 50Mha, which is 5% of total soil C in 0.3% of the world total land area. Indonesia is the country where more than 80% of tropical peat distributes (Maltby and Immirzi 1993). In the peatlands, exploitation has been conducted with clear cutting of the forests and constructing drainage channels. Drainage of the peatlands generally induces peat subsidence with peat decomposition (Couwenberg *et al.*, 2010), and peat decomposition accounts for 50 - 70% of the subsidence (Murayama and Bakar, 1996 a, b; Hatano *et al.*, 2010). Hatano *et al.* (2010)

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estimated that peat decomposition per unit peat subsidence ranged from 4.3tC ha<sup>-1</sup> cm<sup>-1</sup> in a drained natural forest to 6.8 tC ha<sup>-1</sup> cm<sup>-1</sup> in a burnt forest in Palangkaraya, Central Kalimantan, Indonesia. Couwenberg et al., (2010) estimated annual rate of peat decomposition per 10 cm of additional drainage depth as 2.5tC ha<sup>-1</sup>yr<sup>-1</sup> in maximum. Peat decomposition induces the emissions of GHGs (CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O), due to release of huge amount of C and N contained in peat (Melling et al., 2005a, b; Melling et al., 2007, Jauhiainen et al., 2005; Takakai et al., 2006). Plots of CO<sub>2</sub> emission from soil surface against average water table depth in tropical peatlands showed a tendency that the CO<sub>2</sub> emission increased with a drop of water table depth, although the relationship was not significant (Hooijer et al., 2006). Based on the tendency, CO<sub>2</sub> emission from peat decomposition in tropical peatland in South Asia was estimated as 170 MtC  $y^{-1}$ , which is half of the Japanese emission (Hooijer *et al.*, 2006). Emissions of N<sub>2</sub>O and CH<sub>4</sub> from tropical peatlands have not been yet estimated widely, however, as N<sub>2</sub>O production is exponentially correlated to CO<sub>2</sub> production (Mu et al., 2009), the contribution of N<sub>2</sub>O emission can be increased significantly by increase of peat decomposition. CH<sub>4</sub> emission from tropical peatlands relatively low (Jauhiainen et al., 2005; Melling et al., 2005b) compared with the CH<sub>4</sub> emissions from boreal and temperate peatlands (Couwenberg *et al.*, 2010), rather CH<sub>4</sub> is absorbed by peat with the water table depth deeper than 50 cm (Watanabe et al., 2009). But, as previously shown, drop of water table depth increases CO, and N<sub>2</sub>O emissions. Therefore, more study for the evaluation about GHGs emissions associated with land uses of tropical peatland is required in order to create proper managements reducing GHGs emissions in tropical peatlands. In this study, we investigated the relationship between GHGs emissions and soil environmental factors in various kinds of land use in tropical peatlands of Indonesia.

#### **MATERIALS AND METHODS**

Four land uses, cropland, burnt forest, drained natural forest and plantation, were investigated in this study. Three sites of cropland area, three sites of burnt forest area, one site in drained natural forest area were chosen in Kalampangan near Palangkaraya city in Central Kalimantan, and six sites in acacia plantation area in Langgam near Kerinci city in Riau. Peat subsidence,  $CO_2$ ,  $N_2O$  and  $CH_4$  fluxes from the soil and soil environmental factors (water table depth, soil temperature, soil moisture content, soil pH, soil electrical conductivity (EC), soil ammonium ( $NH_4^+$ -N) and nitrate ( $NO_3^-$ -N) contents, water soluble soil organic matter content (WSOC)) were measured once a month from September 2008 to August 2009. Total amount of precipitation during the period of the investigation was similar in Palangkaraya (2146 mm) and Kerinci (2187 mm), but monthly pattern of the precipitation was different. Rainy season, which was defined as the period more than two months with monthly precipitation more than 200, was found from November to April in Palangkaraya, but in Kerinci, rainy season was divided into two periods from September to November and from March to April. Bulk density of top 0-10 cm was 0.38 - 0.42 g cm<sup>-3</sup> in acacia plantation.

CO<sub>2</sub>, N<sub>2</sub>O and CH<sub>4</sub> fluxes were measured by using closed chamber method. Above ground green parts were cut and eliminated before the measurements. Peat subsidence was measured by iron pipe inserted to under-layered mineral soil or PVC pipe inserted to 2 m depth. Water

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table depth was measured by 2 m long perforated PVC pipes. Soil temperature at 4 cm depth was measured by thermometer, and water temperature was measured during flooding instead of soil temperature. Volumetric soil moisture content in 0 to 6 cm depth was measured by TDR moisture meter. Soil pH and soil EC was measured using suspension of 1:20 peat and water mixture. Soil NH<sub>4</sub><sup>+</sup>-N content was measured in the extraction of 1:20 peat and 2M-KCl solution mixture. Soil NO<sub>3</sub><sup>-</sup>-N and WSOC contents were measured in the extraction of 1:20 peat and water mixture.

### **RESULTS AND DISCUSSION**

Monthly change in soil temperature, water table depth, soil moisture content in cropland, natural forest and burnt forest, and acacia plantation are shown in Figure 1. Soil temperature was higher in cropland and burnt forest than in natural forest and acacia plantation, due to open canopy of the vegetation in cropland and burnt forest. Water table depth was deeper in cropland than in other land uses. In burnt forest, flooding was shown in the rainy season. Soil moisture content was high in cropland and low in acacia plantation compared with that in forest and burnt forest.

Monthly change in CO<sub>2</sub>, N<sub>2</sub>O and CH<sub>4</sub> fluxes in cropland, forest and burnt forest and acacia plantation are shown in Figure 2. CO<sub>2</sub> flux was high in cropland and low in forest and burnt forest, compared with that in plantation. Melling *et al.* (2005) showed CO<sub>2</sub> flux was higher in tropical peatland with lower humidity due to open canopy, higher bulk density and deeper water table depth, this explained well the results obtained in this study. However, there were several peaks of CO<sub>2</sub> flux in rainy season. This may indicate that microbial activities were influenced by water stress in dry season (Couwenberg *et al.*, 2010). N<sub>2</sub>O flux was extremely high in cropland, which was about 40 times large compared with that in other land uses, as previously shown (Takakai *et al.*, 2006). There were peaks of N<sub>2</sub>O flux in rainy season in all land uses. CH<sub>4</sub> flux was absorbed in dry seasons, but high in rainy season when water table depth became shallow, as shown in previous reports (Jauhiainen *et al.*, 2005; Melling *et al.*, 2005b; Watanabe *et al.*, 2009). Especially in burnt forest where flooding happened, the CH<sub>4</sub> flux was about 10 times large compared with that in other land uses, but it was 3800 µgC m<sup>-2</sup> hr<sup>-1</sup> in maximum which was low compared with that from the boreal and temperate peatlands (Couwenberg *et al.*, 2010).

Figure 3 shows that the relationship between average  $CO_2$  flux and average soil temperature, average soil moisture content and average water table depth.  $CO_2$  flux was significantly correlated with soil moisture content negatively and with water table depth positively, but not significantly correlated with soil temperature. Correlation coefficient in the relationship between  $CO_2$  flux and water table depth was higher than soil moisture content.

Figure 4 shows the relationship between average water table depth and average  $N_2O$  flux, and average  $CH_4$  flux. Average  $N_2O$  flux was logarithmically correlated to water table depth. This indicates that  $N_2O$  emission is more sensitive to change in water table depth than  $CO_2$  emission, as shown previously (Mu *et al.*, 2009).  $CH_4$  flux was negatively correlated with water table depth, which is the same tendency as previous report (Watanabe *et al.*, 2009; Couwenberg *et al.*, 2010).

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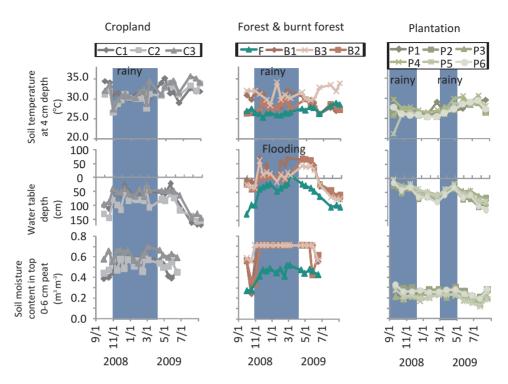
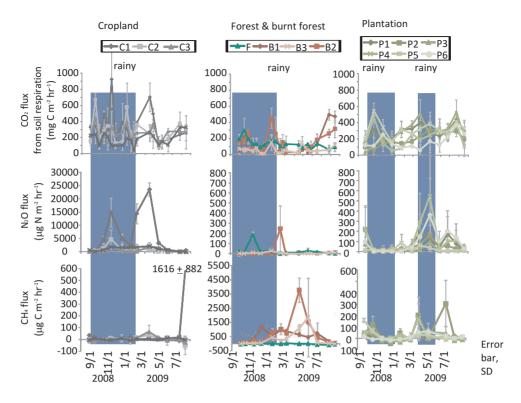


Figure 1. Monthly change in soil temperature, water table depth and soil moisture.

Results of regression analysis between the gas fluxes and soil environmental variables are shown in Table 1. Water table depth was most important factor for controlling GHGs emissions. In addition, soil EC and soil  $NO_3$ -N content were well correlated with the GHGs emissions. Soil  $NO_3$ -N content was highly correlated with N<sub>2</sub>O flux. Taking into consideration peaks of N<sub>2</sub>O flux in rainy season, denitrification is thought to be main process for N<sub>2</sub>O emission. Increase of soil EC indicates an increase of soil nutrient concentrations, which is associated with organic matter decomposition, N mineralization and nitrification. This leads to denitrification consequently. Takakai *et al.*, (2006) showed more N<sub>2</sub>O emission in the year with higher precipitation, and Hashidoko *et al.* (2008) showed the development of a unique denitrifier in  $NO_3$ -N rich peat soil.

 $CH_4$  flux, which increased during a period of flooding, was significantly correlated to WSOC content. Increase of WSOC indicates a reduction of oxidative decomposition of soil organic matter due to decrease of  $O_2$  gas supply into soil from the atmosphere by flooding. Consequently,  $CH_4$ , which was produced from  $Co_2$  and acetate in soil, can be emitted from the soil surface without  $CH_4$  oxidation (Conrad, 1999). However, increase of WSOC also indicates excess  $CO_2$  production due to act of soil organic acid as electron acceptor (Yao and Conrad, 2000).

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Figure 2. Monthly change in CO<sub>2</sub>, N<sub>2</sub>O and CH<sub>4</sub> fluxes.

	CO <sub>2</sub>		log(N <sub>2</sub> O)	)	CH <sub>4</sub>	
Soil temperature	0.009		0.279		0.123	
Water table depth	0.784	**	0.863	**	-0.828	**
Soil moisture	-0.592	*	-0.359		0.666	
Subsidence	0.366		0.343		-0.236	
pН	0.186		0.423		-0.041	
EC	0.689	**	0.590	*	-0.555	*
WSOC	-0.360		-0.423		0.544	*
NH4 <sup>+</sup> -N	0.029		-0.205		-0.283	
NO <sub>3</sub> -N	0.671	*	0.715	**	-0.407	

 Table 1. Results of regression analysis between the gas fluxes and soil environmental variables

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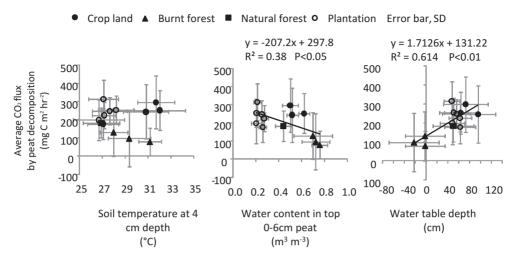


Figure 3. Relationship between average peat decomposition and average soil environmental factors.

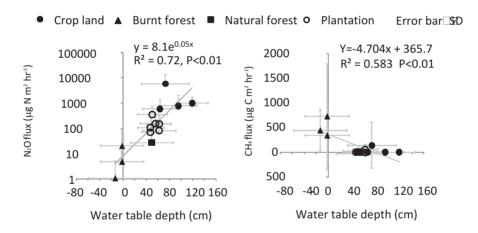


Figure 4. Relationship between water table depth and the fluxes of  $N_2O$  and  $CH_4$ .

Peat subsidence was not correlated with the gas fluxes. However, increase of subsidence increased  $CO_2$  and  $N_2O$  fluxes and decreased  $CH_4$  flux. Peat subsidence is influenced by bulk density, and peat subsidence associated to the decomposition of peat with higher bulk density produces more  $CO_2$ .

Figure 5 shows the relationship between global warming potential (GWP) and water table depth. The GWP of  $CO_2$  base was converted from GHGs emissions using the factors which 25 g  $CO_2$  for 1 g  $CH_4$  and 298 g  $CO_2$  for 1 g  $N_2O$  at the 100-year time horizon (Forster *et al.*, 2007).

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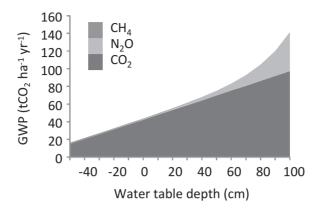


Figure 5. Relationship between water table depth and global warming potential (GWP)

GWP was estimated to increase from 40 to 140 t CO<sub>2</sub> ha<sup>-1</sup> yr<sup>-1</sup>, when water table dropped from 0 to 100 cm. N<sub>2</sub>O emission significantly contributed to the GWP, when water table became deeper than 50 cm, and it occupied 29 % of GWP at water table depth of 100 cm. However, the contribution of CH<sub>4</sub> emission to GWP was small even during the period of flooding.The relationship between CO<sub>2</sub> emission and water table depth given by Hooijer et al. (2006) was a straight line which CO<sub>2</sub> emission increases from 0 to 80 t CO<sub>2</sub> ha<sup>-1</sup> yr<sup>-1</sup> when water table dropped from 0 to 100 cm. Therefore, the result from this study will make considerably (75%) large estimation of peat decomposition compared with Hooijer et al. (2006). Concerning the CO, emission from flooded tropical peatland, several researchers reported the values of the CO, emission from the peatland, where the annual water table depth ranged from -5 to 2 cm and was -0.6 cm in average, ranged from 4.1 to 88 t CO<sub>2</sub> ha<sup>-1</sup> yr<sup>-1</sup> and was 32t CO<sub>2</sub> ha<sup>-1</sup> yr<sup>-1</sup> in average (Chimner, 2004; Chimner & Ewel 2004; Furukawa et al., 2005; Hadi et al., 2001; Jauhiainen et al., 2001; Darung et al., 2005). The average value of CO<sub>2</sub> emission may be similar as net primary production (NPP) of natural tropical forest. Hirano et al. (2007) showed that net ecosystem CO<sub>2</sub> exchange (NEE = soil CO<sub>2</sub> emission - NPP) was 3.7 and 29.0 t CO<sub>2</sub> ha<sup>-1</sup> yr<sup>-1</sup> in a natural swamp forest and a well flooded burnt forest in Palangkaraya, respectively. The difference between the values of NEE in natural and burnt forests (25 t CO<sub>2</sub> ha<sup>-1</sup> yr<sup>-1</sup>) can be an estimation of NPP of trees of natural forest of Palangkaraya. However, Couwenberg et al., (2010) reported that NEE values in temperate peatlands were negative (which means the absorption of  $CO_2$  by the ecosystem) when water table depth existed in 0 to 10 cm, and around - $5 \text{ t CO}_{2} \text{ ha}^{-1} \text{ yr}^{-1}$  at 0 cm of water table depth. These findings indicate that large amount of peat is always decomposed even in natural forest in tropical peatlands. Vegetation fixes CO<sub>2</sub> emitted from soil as well as Co<sub>2</sub> in atmosphere, so the loss of vegetation from the peatlands indicates that CO<sub>2</sub> emitted from soil is converted into the atmosphere. Therefore, in order to reduce NEE, proper plantation to increase NPP is also important as well as the decrease of CO<sub>2</sub> and N<sub>2</sub>O emissions from peatland by the maintenance of water table depth.

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#### CONCLUSION

Water table depth was a significant factor for determining GHGs emissions from tropical peatlands.  $CO_2$  emission from the peatlands increased from 40 to 100 t  $CO_2$  ha<sup>-1</sup> yr<sup>-1</sup> with drop of water table from 0 to 100 cm. Increase of peat decomposition promoted N<sub>2</sub>O emission due to increase of N mineralization and nitrification followed by denitrification in wet season. N<sub>2</sub>O emission significantly contributed to GWP at the water table depth deeper than 50 cm.  $CH_4$  emission increased with flooding, but it was not significant in GWP.

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